

Paleotempestology of Onslow Bay, NC

Scott Hippensteel

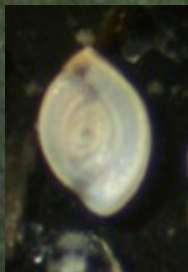
University of North Carolina at Charlotte



Methodology/Goal

Findings

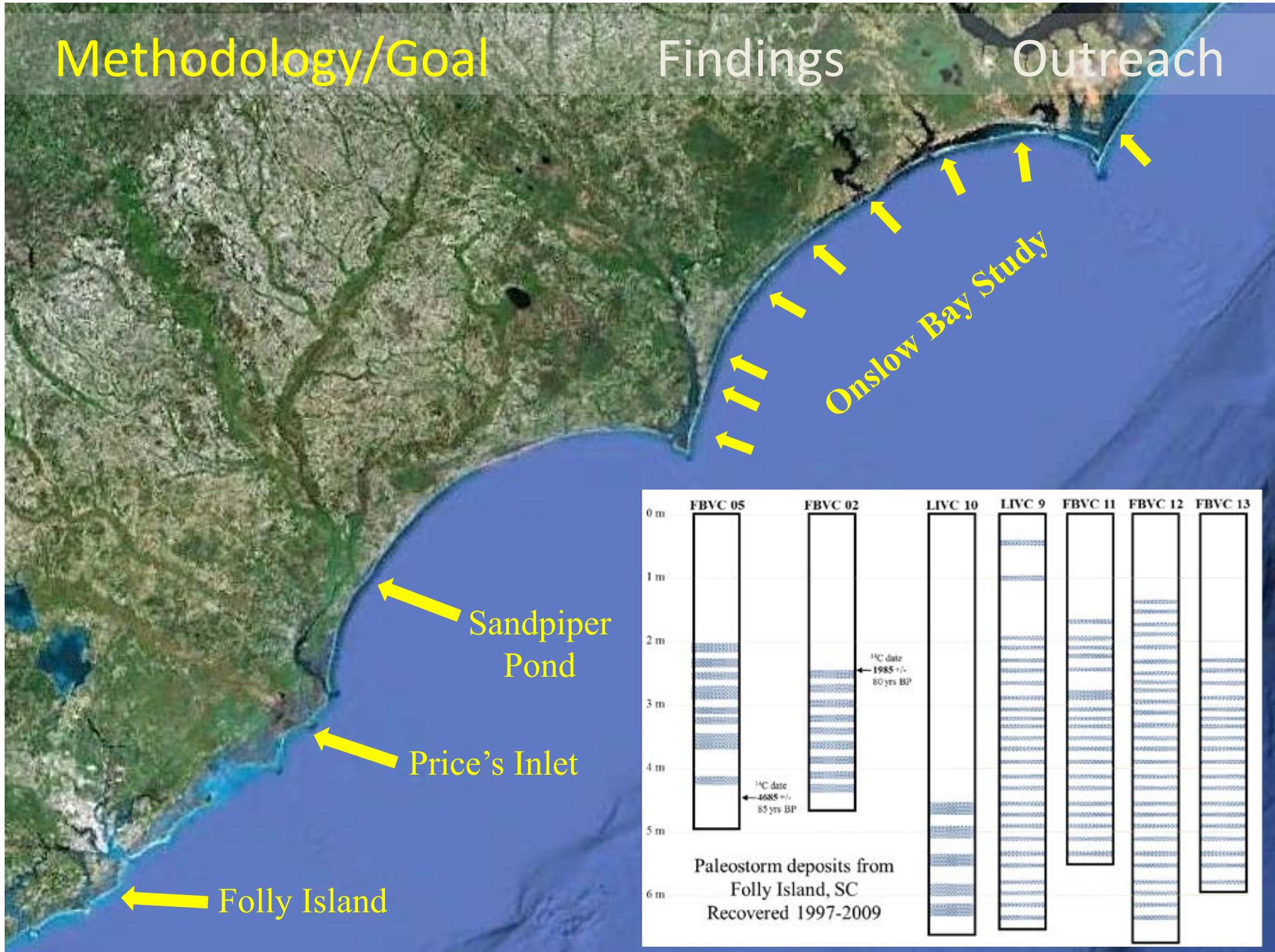
Outreach



Methodology/Goal

Findings

Outreach



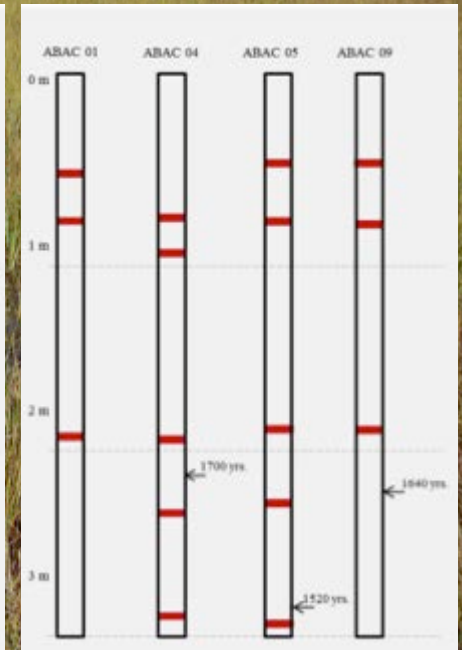
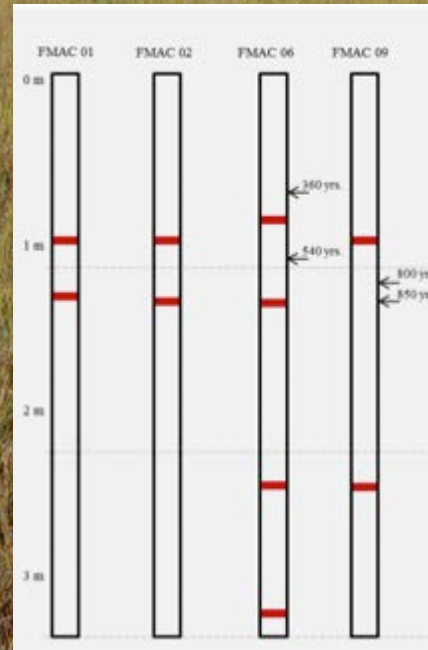
Methodology/Goal

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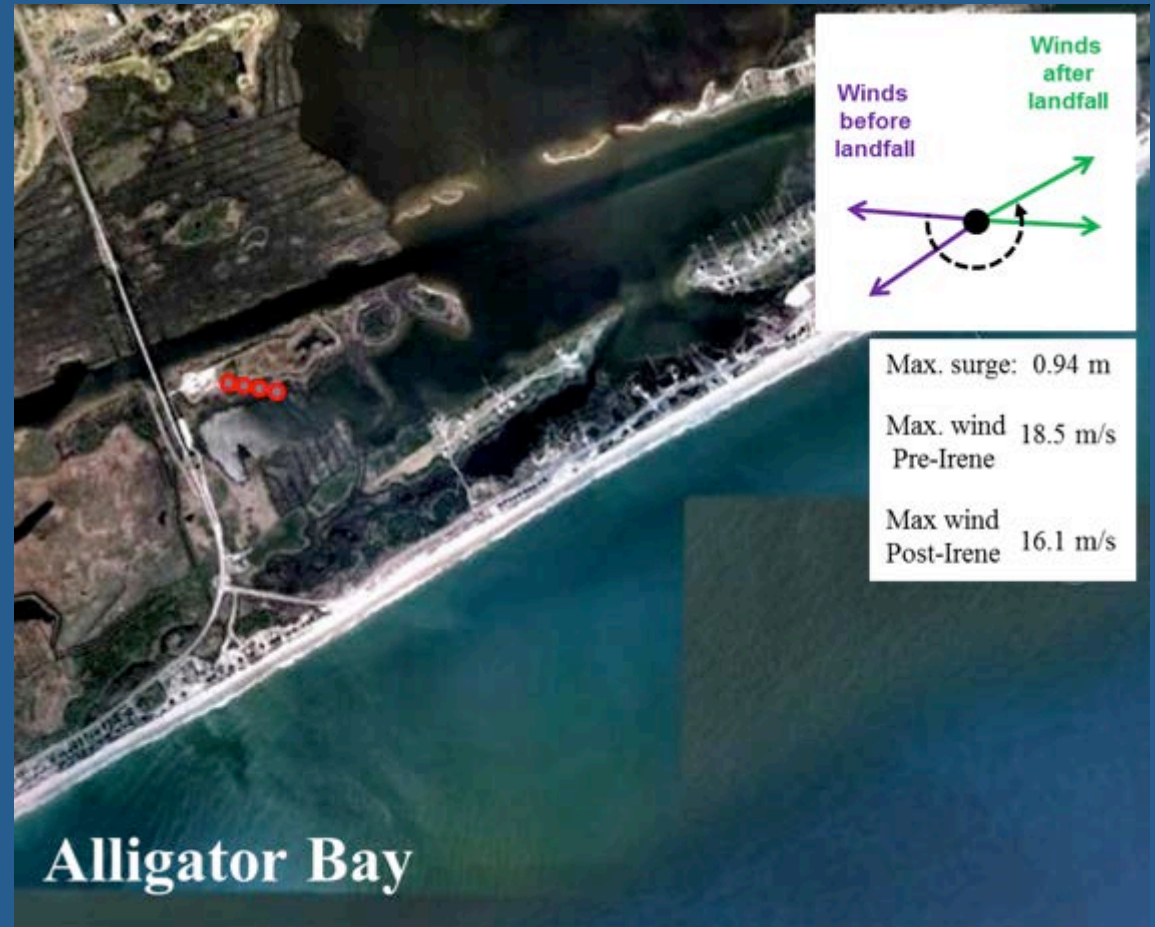
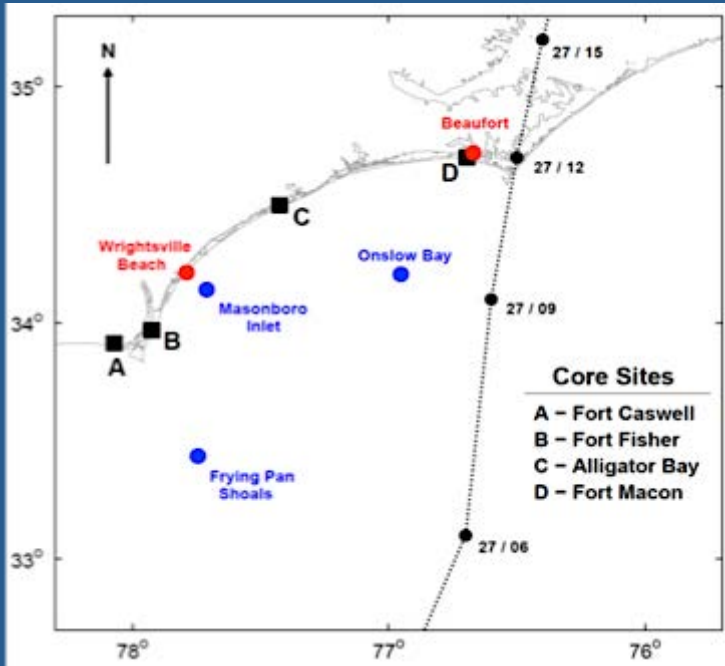
	Composite (total) <u>Storm Deposits</u>	<u>Storms</u>
- South Core Banks	0	0
- Tar Landing Bay	10	5-10 (?)
- Hope Pole Creek	0	0
- Bear Island	0	0
- Alligator Bay	16	6-16 (?)
- Wrightsville Beach	0	0
- Fort Fisher	0	0
- Battery Buchanan	0	0
- Oak Island	0	0



Methodology/Goal

Findings

Outreach



Methodology/Goal

Findings

Outreach



Coastal Geology
Applied Paleontology
Environmental Micropaleontology

Methodology/Goal

Findings

Outreach



Paleotempestology and the pursuit of the perfect paleostorm proxy

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Growing populations and recent hurricane activity along the Atlantic and Gulf coasts have made clear the need for a more accurate and extensive record of storm activity. The paucity of historical descriptions of hurricane strikes along the eastern Atlantic coast limits their use as a predictive tool. Only three Category 5 hurricanes have made landfall in the United States since 1900, as a result, the likelihood of a major storm impacting the region cannot be accurately determined using historical records. Proxy records collected from marginal marine environments offer the potential to extend the record back several thousand years, providing better statistical confidence on hurricane problems and a better understanding of the influence of global warming on catastrophic hurricane development.

Although definitive paleotempest deposits from North America date back to the Oligocene, "paleotempestology" is a relatively new field of research (Dyer and Dyer, 1964). Two proxy approaches developed independently in the early 1980s from two different geographic regions (the resulting storm records offer increasing insight into the problems associated with the existing methodology. At the same time that a research team from Louisiana State University was using sedimentary criteria to detect storm deposits in the coastal lakes of Louisiana, I was finding offshore indicator fossilization (and Oligo-Miocene land) in buried sand layers from the back-barrier marshes of South Carolina (Lu and Frazer, 1993, 2001; Hopper and Martin, 1995, 1996, 2000; Lu, 1996, 2003). Over the next decade, several follow-up studies from the same and different regions led to a series of advances in the validity of each approach (Donnelly et al., 2001, 2003a, 2003b, 2004; Scott et al., 2005; Donnelly, 2005; Hopper et al., 2005; Lambert et al., 2006; Lu et al., 2006).

Precious well-dated paleotempest proxies have been hindered by two primary issues. First, studies based on sand layers in marshes and coastal lakes have been criticized because the mechanism of deposition was not always well-documented, and the source of the sediment was not clearly defined (Dronin, 1999, 2002, 2006; Lu et al., 2005). Second, the depositional history of many marsh records is short, especially in New England and the mid-Atlantic (Donnelly et al., 2005; Tink, 2002). The use of microfossil-enriched storm layers from near

Charleston, South Carolina, avoided both limitations. The mechanism for deposition is well understood because the storm deposits contain offshore indicator fossilization, and the record extends over at least the mid-Holocene (Hopper and Martin, 1995, 2003).

The micropaleontological proxy is not without problems or critics, however. The sand layers included in a list of "offshore indicators" genera and the role of bioturbation in the preservation of storm deposits has been strongly debated (e.g., Scott et al., 2005, 2007; Hopper et al., 2005; Hopper and Martin, 2000). Further, the geochronology established for many of these studies is questionable. The weathered deposits themselves are, by their very nature, the product of severe reworking of ancient offshore sediments. Dating the fossilization in such deposits directly would yield the age of the fossils (Oligo-Miocene), not the time of storm deposition.


As a result, the quality, completeness, and comparability of the storm record from the southeastern Atlantic is not known. Better records are dynamic environments, which, in a management setting, will cover on themselves, destroying or obscuring the back-barrier storm record. Additionally, marsh subsidence in a progressive setting will shift through time. In South Carolina, for example, paleostorm frequency appears to be decreasing through time, with more storms lower in the marsh strata (Hopper and Martin, 2006; Hopper et al., 2006). Is this apparent decrease in catastrophic storm events a result of climate change or an artifact of the midline processes of bioturbation and reworking? Is taphonomy partly to blame for the potentially poor preservation, or even absence, of some microfossils in the storm deposits? These marsh subsidence issues have, after all, been flooded from high marsh to low marsh, and the primary back-barrier process, higher water, prefer the low marsh. Before paleotempest proxies reach their potential as a tool for paleoclimatic researchers, their geological and biological factors need to be better understood.

The field of paleotempestology has never been of more importance. Paleohistorical storm data would be more valuable in the current period of climate change, despite the frustratingly elusive nature of the "ideal" paleotempest record. This ideal paleotempest record, whenever and wherever it is discovered, will have four interrelated characteristics:


1. The deposits will display a clear link to the source of sediments and the mode of deposition. This may be because of the fossilization present in the deposit, some other

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The geological legacy of Hurricane Irene: Implications for the fidelity of the paleo-storm record